Water Samples and Evaporation Estimates Interim Report

C-15877, Biscayne Bay Paleoecological Salinity Study
January 26, 2004

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Task 2

This Task 2 report represents a collaboration between Rene Price at Florida International University and Peter Swart. The section on evaporation rates calculated from insolation data was contributed by Rene Price.

Water Sample Data

Water samples collected by the Southeast Environmental Research Center (SERC) at Florida International University (FIU) have been analyzed for their hydrogen, oxygen, and carbon isotopic composition. Samples have been analyzed through October 2003. Sample locations are given shown in Figure 1 and all data are shown in Appendix 2. Graphs of the temporal changes in the oxygen isotopic composition for the 24 stations in Biscayne Bay are contained in Appendix 1. In addition samples from several months have been analyzed for their minor elemental composition. These data are shown in Appendix 3.

Figure 1: Sample location from which water samples are collected at monthly intervals by SERC. These are being analyzed for stable oxygen, hydrogen, and carbon isotopes in this study.

Status of Water Sampling

Water samples are routinely collected by SERC and analyzed by the stable isotope laboratory at the University of Miami. All samples collected to date have been analyzed for O and H isotopes

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Figure 2: Evaporation estimates for Biscayne Bay.

with the exception of November and December 2003 and January 2004. Occasionally water samples are not sub-sampled for stable isotopes. We are working to establish a better system to ensure all samples are collected for us in a reliable manner.

**Evaporation Rates calculated from Insolation Data**

Climatic data is not yet available from the station on Elliot Key, therefore, evaporation rates using direct data from Biscayne Bay could not be determined. The nearest weather station is located in Key Largo on the boat dock at the Ranger Station of Everglades National Park. This station was installed by our group and is currently being maintained without direct support. Climatic data available from this station includes, air temperature, relative humidity, wind speed and direction, incoming radiation, and rain amount. These data were collected at 30-minute intervals for an approximately 2-year period between September 2001 and September 2003. Evaporation estimates were obtained using the incoming radiation and air temperature data from Key Largo and applying the data to the Priestly-Taylor formula for evaporation. The Priestly-Taylor formula uses an energy budget method for estimating evaporation and is based on the
Table 1. Summary of monthly evaporation rates at the Key Largo Ranger Station.

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* data missing from multiple days during these months

Principal that the flux of heat associated with evaporation, i.e. the latent heat flux, is a major flux in the energy budget of the evaporating surface. Principal fluxes in the energy budget at the Earth's surface are of two types, radiation fluxes (long-wave and short wave) and advective fluxes (evaporation and conduction). The physics of energy exchange with laterally extensive, wet surfaces is such that both evaporation and conduction vary with the net radiation flux to the surface (Brutsaert 1982). The Priestly-Taylor formula uses the equation:

\[ E = \alpha \frac{\Delta}{\Delta + \gamma} R_{NET} \]  \hspace{1cm} (1)
in which, RNET is the net radiation flux at the surface (W/m²), \( a \) is the slope of the saturation water vapor pressure curve at the temperature of the air, and \( \alpha \) is the psychrometric constant. Only incoming (short-wave) radiation was measured at the Key Largo station. The short-wave radiation (Rshort) measured at our station on Key Largo was converted to a net radiation flux (RNET) using a correlation relationship obtained from a climatic station located near Butternut Key in Florida Bay. This relationship was determined to be \( \text{RNET} = 6.113 \text{(Rshort)} - 44.89 \) with and \( R^2 = 0.988 \). A value for the empirical coefficient in the Priestly-Taylor formula of \( \beta = 1.26 \) has been found to work well over a large range of conditions, particularly in large wet areas such as the Everglades (German 2000), Florida Bay and Biscayne Bay (German 2000), and this value was used in this exercise.

The results of the evaporation estimates are summarized in Table 1 and illustrated in Figure 2. Evaporation rates varied seasonally. Monthly evaporation rates were lowest in December (9.4 cm in both 2001 and 2002), and highest in the spring and summer months of April through July (17 - 22.4 cm per month). Gaps in the data during several months of the two-year monitoring period prevented an annual estimate in evaporation. However, annual estimates of evaporation in Florida Bay range from 125 to 169 cm/year (Nuttle et al., 2003).

Although the evaporation estimates were made based upon radiation and air temperature data from Key Largo, similar conditions are expected in Biscayne Bay. Once climatic data becomes available from Elliot Key, we will be able to make estimates of evaporation more precisely for Biscayne National Park.

The preceding section was contributed by Rene Price at Florida International University.

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Evaporation Rates from Stable Isotopic Data

The stable O and H isotopic composition of waters increase as a result of evaporation. This increase is the corner stone of the use of the stable isotopic ratio in corals, ostracods, and foraminifera for paleoenvironmental analysis not only in this project, but also in global studies. Generally speaking fossils which have higher oxygen isotopic compositions are derived from environments with higher salinities. Hence the oxygen and hydrogen isotopic composition of the water is an indicator of salinity. The extent of the increase in the oxygen isotopic composition however is dependent not only on salinity, but also upon the temperature of evaporation, the relative humidity of the atmosphere, and the isotopic composition of the atmosphere. Figures 3 and 4 show the increase in the oxygen isotopic composition of an evaporating basin such as Biscayne Bay (Figure 3) and a freshwater body of water such as the Everglades (Figure 4). The different initial isotopic compositions reflect that fact that the origin of the water in the Everglades is local precipitation which has a oxygen isotopic composition of -3 ‰, while salt water which contributes to Biscayne Bay is derived from the Gulf Stream and has an isotopic composition of approximately 0.8 to 1‰. The different behavior of the oxygen isotopes at higher percentages of evaporation is a result of changes in the activity coefficients of water in solution of high ionic strength. In figure 3 and

Figure 5: Theoretical change in the oxygen isotopic composition and salinity of Biscayne Bay as a result of evaporation.

Figure 6: Salinity and oxygen isotopic data from BNP superimposed on a evaporation model for BNP.

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4 the x-ordinate can be converted to a salinity scale assuming that the initial water body has a specific salinity. For example in Figure 5, the initial salinity is taken as being the same as Gulf Stream Seawater. If this water is allowed to penetrate into Biscayne Bay, then its isotopic composition would behave according to figure 5. As may be observed the maximum oxygen isotopic composition which could be attained in an atmosphere of 75% relative humidity (similar to the average observed in South Florida) would be +4‰ if the salinity reached as high as 50. The salinity and oxygen isotopic data can be plotted of a figure such as that shown in Figure 4 to obtain an idea of the amount of evaporation. This is shown in Figure 7. However, in order to utilize the stable isotopic data to estimate amounts of evaporation, a mass balance of the inputs of water flowing in and out of Biscayne Bay must be constructed. Such a mass balance can be described by the following equation.

$$Q_{I1} + Q_{I2} + Q_{I3} + Q_p = Q_{O1} + Q_E$$

In this equation Q represents the flux of water in (I) and out (O) of Biscayne Bay. Specifically:
- $Q_{I1}$ = flux derived from ocean
- $Q_{I2}$ = flux from land
- $Q_{I3}$ = flux from groundwater
- $Q_{O1}$ = flux to the ocean
- $Q_p$ = flux from precipitation
- $Q_E$ = flux from evaporation

Similar equations can be constructed for salinity and isotopic composition. As salinity of freshwater = 0 then the salinity equation can be reduced to:

$$Q_{IS2} + Q_{IS3} = Q_{OS1}$$

For stable oxygen isotopes, the equation becomes:

$$Q_1 \delta_1 + Q_1 \delta_2 + Q_1 \delta_3 + Q_p \delta = Q_{O1} \delta_1 + Q_E \delta_E$$
At the present time the principal unknowns in these equations are the fluxes from the oceans, the flux from groundwater, and the evaporative flux. At the present time we are collecting information of freshwater discharge and precipitation. Estimate of evaporation presented in this report can be utilized with an estimate of groundwater input to determine the oceanic flux. Alternatively the oceanic flux can be derived from various hydrodynamic models which are being constructed.

**Figure 7: Conceptual hydrodynamic model for Biscayne Bay.**

- $Q_P$ and $Q_E$ are the oceanic fluxes, $Q_I$ for freshwater discharge, $Q_O$ for precipitation, and $Q_I^2$ for groundwater input.

**Variances from Projected Work Tasks and/or Schedule**

All items of the proposed work are on track. The only exception is that the weather station due to be installed by BNP has not yet been installed. The installation of this station is beyond our control. The sensor for the station which was necessary for the measurement of insolation was purchased by the University of Miami and delivered to BNP over 12 months ago. I will agitate that this be carried out as soon as possible. If the successful installation is delayed then we will be able to utilize a combination of insolation data from Florida Bay and from the weather stations at RSMAS for the final tuning of the evaporation model.

**References**


Appendix 1: Graphs of temporal variation of oxygen isotopic composition in Biscayne National Park
Appendix 2: Data Tables of oxygen, hydrogen, and carbon isotopic composition of surface waters in Biscayne National Park
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| Min          | 0.8 | 1.6 | 0.8 | -0.2 | 0.7 | 0.1 | 0.3 |
| Std          | 0.2 | 0.3 | 0.4 | 0.4 | 0.3 | 0.3 | 0.3 |
|      | BISC       | Convoy Point | 103 BISC Near Black Ledge | 104 BISC BNP Marker C | 108 BISC Marker G-1B | 109 BISC Midbay North | 110 BISC Fender Point | 111 BISC Featherbed Bank | 112 BISC Sands Cut | 113 BISC Elliott Key | 116 BISC Rubicon Keys | 121 BISC Card Sound North | 122 BISC West Arsenicker | 123 BISC Pelican Bank | 124 BISC Midbay South | 126 BISC BNP Marker B | 127 BISC Shoal Point | 128 BISC Matheson Beach | 129 BISC Marker G-71 | 130 BISC South Dodge Island | 131 BISC North Venetian Basin | 132 BISC North I-195 Basin | 133 BISC North Normandy Isle | 134 BISC Oleta River Park | 135 BISC South Card Sound |
|------|------------|--------------|---------------------------|----------------------|---------------------|----------------------|----------------------|------------------------|---------------------|----------------------|------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|
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| 86   | Apr-98 | 1.4 | 1.6 | 1.0 | 1.6 | 1.8 | 1.6 | 1.6 | 1.4 | 1.8 | 1.6 | 1.6 | 1.4 | 1.8 | 1.6 | 1.6 | 1.4 | 1.8 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 |
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| 88   | Jun-98 | 2.9 | 2.6 | 2.0 | 2.6 | 2.5 | 2.0 | 2.9 | 2.0 | 2.9 | 2.0 | 2.9 | 2.0 | 2.9 | 2.0 | 2.9 | 2.0 | 2.9 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 |
| 89   | Jul-98 | 0.9 | 1.2 | 0.7 | 1.2 | 0.9 | 1.2 | 1.0 | 0.7 | 1.0 | 0.7 | 1.0 | 0.7 | 1.0 | 0.7 | 1.0 | 0.7 | 1.0 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 |
| 90   | Aug-98 | 0.5 | 0.8 | 0.4 | 0.8 | 0.8 | 0.4 | 0.5 | 0.4 | 0.5 | 0.4 | 0.5 | 0.4 | 0.5 | 0.4 | 0.5 | 0.4 | 0.5 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 |
| 91   | Sep-98 | 0.7 | 1.1 | 0.6 | 1.1 | 1.1 | 0.6 | 0.7 | 0.6 | 0.7 | 0.6 | 0.7 | 0.6 | 0.7 | 0.6 | 0.7 | 0.6 | 0.7 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 | 0.6 |
| 92   | Oct-98 | 1.3 | 1.7 | 0.9 | 1.3 | 1.7 | 0.9 | 1.3 | 1.7 | 0.9 | 1.3 | 1.7 | 0.9 | 1.3 | 1.7 | 0.9 | 1.3 | 1.7 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 |
| 93   | Nov-98 | 0.9 | 1.3 | 0.7 | 0.9 | 1.3 | 0.7 | 0.9 | 1.3 | 0.7 | 0.9 | 1.3 | 0.7 | 0.9 | 1.3 | 0.7 | 0.9 | 1.3 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 |
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**Biscayne Bay Oxygen**

- **Biscayne Bay Oxygen**
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- **103 BISC Near Black Ledge**
- **104 BISC BNP Marker C**
- **108 BISC Marker G-1B**
- **109 BISC Midbay North**
- **110 BISC Fender Point**
- **111 BISC Featherbed Bank**
- **112 BISC Sands Cut**
- **113 BISC Elliott Key**
- **116 BISC Rubicon Keys**
- **121 BISC Card Sound North**
- **122 BISC West Arsenicker**
- **123 BISC Pelican Bank**
- **124 BISC Midbay South**
- **126 BISC BNP Marker B**
- **127 BISC Shoal Point**
- **128 BISC Matheson Beach**
- **129 BISC Marker G-71**
- **130 BISC South Dodge Island**
- **131 BISC North Venetian Basin**
- **132 BISC North I-195 Basin**
- **133 BISC North Normandy Isle**
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- **135 BISC South Card Sound**
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### Summary Statistics

- **Mean**: 4.8, 18.0, 14.0, 7.7, 7.0, 6.8, 5.3, 9.6
- **Maximum**: 14.3, 18.6, 19.3, 19.9, 10.7, 9.7, 10.2, 17.6
- **Minimum**: 0.2, 17.3, 9.3, -4.3, -0.6, 3.7, -1.8, 0.0
- **Standard Deviation**: 4.3, 0.7, 3.0, 5.9, 3.1, 1.4, 3.3, 4.8
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| Minimum                       | 5.0    | -3.7   | -2.9   | 5.9    | 5.1    | -1.6   | 7.8    | 9.8    | -1.4   | 3.2    | 4.0    |
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**Maximum**: 31.0 16.1 8.6 13.7 13.1 12.3 17.0

**Minimum**: 10.1 4.5 -1.6 1.8 0.3 -2.7 -17.7

**Standard Deviation**: 6.3 3.7 3.2 3.0 3.4 4.0 6.8
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